Chemistry 2000 Slide Set 11: Chemical equilibrium

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Thermodynamic criterion for equilibrium

 Recall that under given reaction conditions (concentrations of reactants and products),

$$\Delta_r G_m = \Delta_r G_m^\circ + RT \ln Q$$

- If $\Delta_r G_m < 0$, the reaction is thermodynamically allowed as written.
- If $\Delta_r G_m > 0$, the reverse of the reaction is thermodynamically allowed.
- What happens if $\Delta_r G_m = 0$?
 - Neither the forward nor reverse direction of the reaction is thermodynamically allowed, so the reaction is at equilibrium.

The equilibrium constant

• At equilibrium,

$$\Delta_r G_m = \Delta_r G_m^\circ + RT \ln Q = 0$$

$$\therefore \Delta_r G_m^\circ = -RT \ln Q$$

- This last equation implies that there is a fixed value of the reaction quotient *Q* when we reach equilibrium.
- We call this value the equilibrium constant *K*. Thus,

$$\Delta_r G_m^\circ = -RT \ln K$$

- For a given reaction, K is a number that depends only on the temperature.
- At equilibrium, Q = K.

• If we know *K*, we can calculate the standard free energy change for a reaction by

$$\Delta_r G_m^\circ = -RT \ln K$$

• We can also rearrange this equation to calculate K from $\Delta_r G_m^{\circ}$:

$$\ln K = -\frac{\Delta_r G_m^{\circ}}{RT}$$

$$\therefore K = \exp\left(-\frac{\Delta_r G_m^{\circ}}{RT}\right)$$

 $[\exp(x) = e^x]$

 $\Delta_r G_m^\circ = -RT \ln K$

Important: K is related to the standard free energy change $\Delta_r G_m^{\circ}$ (obtained from tables of standard free energies of formation), not to $\Delta_r G_m$, which is zero at equilibrium. Thinking in terms of Q and K

• Roughly, $Q = \frac{\text{products}}{\text{reactants}}$.

- If Q < K, then Q needs to grow to reach equilibrium.
- Q increases if we make more products and consume reactants.

Conclusion: If Q < K, a reaction is thermodynamically allowed as written.

Converse: If Q > K, the reverse of the reaction is thermodynamically allowed.

Conditions for thermodynamic feasibility

Example: Dimerization of NO₂

$$2NO_{2(g)} \rightleftharpoons N_2O_{4(g)}$$

$$\Delta_r G_m^\circ = \Delta_f G^\circ(N_2O_4, g) - 2\Delta_f G^\circ(NO_2, g)$$

= 97.79 - 2(51.3) kJ/mol = -4.8 kJ/mol
$$\mathcal{K} = \exp\left(-\frac{\Delta_r G_m^\circ}{RT}\right)$$

= $\exp\left(-\frac{-4.8 \times 10^3 \text{ J/mol}}{(8.314463 \text{ J K}^{-1} \text{ mol}^{-1})(298.15 \text{ K})}\right)$
= 6.96

(Note the double negative in the exponential.)

- Suppose that, at some particular point in time, $p_{\rm NO_2} = 0.4$ bar and $p_{\rm N_2O_4} = 1.8$ bar.
- Is the reaction thermodynamically allowed under these conditions?

$$Q = \frac{a_{N_2O_4}}{(a_{NO_2})^2} \\ = \frac{p_{N_2O_4}/p^{\circ}}{(p_{NO_2}/p^{\circ})^2} \\ = \frac{1.8}{(0.4)^2} = 11 \\ Q > K = 6.96$$

Conclusion: The reaction will run backwards, i.e. N_2O_4 will dissociate into NO_2 .

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• We could have come to the same conclusion by calculating $\Delta_r G_m$:

$$\begin{split} \Delta_r G_m &= \Delta_r G_m^\circ + RT \ln Q \\ &= -4.8 \, \text{kJ/mol} + (8.314\,463 \times 10^{-3} \, \text{kJ} \, \text{K}^{-1} \text{mol}^{-1}) \\ &\times (298.15 \, \text{K}) \ln(11) \\ &= 1.2 \, \text{kJ/mol} > 0, \end{split}$$

from which we also conclude that the reverse reaction is thermodynamically allowed.

Calculating an equilibrium mixture

$$2NO_{2(g)} \rightleftharpoons N_2O_{4(g)}$$

- For the composition given above (0.4 bar NO₂, 1.8 bar N₂O₄, we know that the reaction will run backwards, dissociating N₂O₄, until equilibrium is reached.
- What is the equilibrium composition?

$$2NO_{2(g)} \rightleftharpoons N_2O_{4(g)}$$

• Use an Initial/Change/Equilibrium (ICE) table to figure out the equilibrium composition.

	NO_2	N_2O_4
Ι	0.4	1.8
С	2 <i>x</i>	-x
Е	0.4 + 2x	1.8 – <i>x</i>

• Substitute into the equilibrium relationship:

$$K = 6.96 = \frac{a_{N_2O_4}}{a_{NO_2}^2} = \frac{1.8 - x}{(0.4 + 2x)^2}$$

and solve for x.

• Find x = 0.0507, which gives $p_{NO_2} = 0.50$ bar and $p_{N_2O_4} = 1.75$ bar.

Acid ionization constants

The acid ionization constant K_a of (e.g.) hydrofluoric acid is the equilibrium constant for the reaction

$$\mathsf{HF}_{(\mathsf{aq})} \rightleftharpoons \mathsf{H}^+_{(\mathsf{aq})} + \mathsf{F}^-_{(\mathsf{aq})}, \quad \mathcal{K}_{a} = 6.6 imes 10^{-4}$$

Problem: Calculate
$$\Delta_f G^{\circ}(HF, aq)$$
 given
 $\Delta_f G^{\circ}(F^-_{(aq)}) = -281.52 \text{ kJ/mol.}$
Answer: $\Delta_f G^{\circ}(HF, aq) = -299.7 \text{ kJ/mol}$

Ocean acidification



Important definition: $pH = -\log_{10}(a_{H^+})$

pH of water in equilibrium with atmospheric CO₂

- CO₂ in the atmosphere currently has a concentration of approximately 411 ppm.
- CO₂ reacts with water according to

$$\mathrm{CO}_{2(g)} + \mathrm{H}_2\mathrm{O}_{(I)} \rightleftharpoons \mathrm{HCO}_{3(\mathsf{aq})}^- + \mathrm{H}_{(\mathsf{aq})}^+$$

- Equilibration with CO₂ therefore acidifies water.
- What is the pH of water that has been equilibrated with the atmosphere at 25 °C?

pH of water in equilibrium with atmospheric CO₂

- ppm can refer to mass or mole fraction.
- In the case of atmospheric gases, the concentration is a mole fraction.
- 411 ppm means that for every million molecules in the atmosphere, approximately 411 are CO₂ molecules.
- Given the proportionality between *n* and *p*, this means that

$$ho_{{
m CO}_2}=rac{411}{10^6}(1.013\,25\,{
m bar})=4.16 imes10^{-4}\,{
m bar}$$

at average sea level pressure.

pH of water in equilibrium with atmospheric CO₂

$$\begin{array}{c} {\rm CO}_{2({\rm g})} + {\rm H}_2 {\rm O}_{({\rm I})} \rightleftharpoons {\rm HCO}^-_{3({\rm aq})} + {\rm H}^+_{({\rm aq})} \\ \\ \\ \hline \\ \frac{\Delta_f G^\circ / {\rm kJ\,mol}^{-1}}{{\rm CO}_{2({\rm g})} & -394.37} \\ {\rm HCO}^-_{3({\rm aq})} & -586.8 \\ {\rm H}_2 {\rm O}_{({\rm I})} & -237.140 \end{array}$$

From these data, calculate

$$\Delta_r G^\circ = 44.7 \, \mathrm{kJ \, mol^{-1}}$$

 $K = 1.47 imes 10^{-8}$

pH of water in equilibrium with atmospheric CO_2

$$CO_{2(g)} + H_2O_{(I)} \rightleftharpoons HCO_{3(aq)}^- + H_{(aq)}^+$$
$$K = \frac{(a_{HCO_3^-})(a_{H^+})}{(a_{CO_2})(a_{H_2O})}$$

- Take $a_{\rm H_2O} \approx 1$.
- $a_{\rm CO_2} = p_{\rm CO_2}/p^\circ = 4.16 \times 10^{-4}$
- If we start with pure water, then $a_{HCO_3^-} = a_{H^+}$.
- Solve

$$1.47 \times 10^{-8} = rac{(a_{\mathsf{H}^+})^2}{(4.16 imes 10^{-4})}$$

to get $a_{\rm H^+} = 2.47 \times 10^{-6}$.

• $pH = -\log_{10}a_{H^+} = 5.61$

Ocean acidification

- The foregoing calculation shows that surface water acidity is linked to atmospheric CO₂ levels.
- More $CO_2 \implies$ increased acidity
- Seawater is slightly alkaline due to equilibria involving carbonate minerals, notably

$$CaCO_{3(s)} \rightleftharpoons Ca^{2+}_{(aq)} + CO^{2-}_{3(aq)}.$$

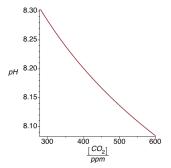
Ocean acidification (continued)

• Key equilibria in sea water

$$\begin{split} \mathsf{CaCO}_{3(\mathsf{s})} &\rightleftharpoons \mathsf{Ca}_{(\mathsf{aq})}^{2+} + \mathsf{CO}_{3(\mathsf{aq})}^{2-} \\ \mathsf{CO}_{3(\mathsf{aq})}^{2-} + \mathsf{H}_2\mathsf{O}_{(\mathsf{I})} &\rightleftharpoons \mathsf{HCO}_{3(\mathsf{aq})}^{-} + \mathsf{OH}_{(\mathsf{aq})}^{-} \\ \mathsf{CO}_{2(\mathsf{g})} + \mathsf{H}_2\mathsf{O}_{(\mathsf{I})} &\rightleftharpoons \mathsf{HCO}_{3(\mathsf{aq})}^{-} + \mathsf{H}_{(\mathsf{aq})}^{+} \end{split}$$

Ocean acidification (continued)

• Calculation of pH vs atmospheric CO₂ pressure using only these equilibria:



(280 ppm = pre-industrial, 411 = current, 600 = year 2100 if emissions continue at current levels)

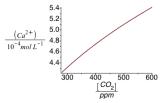
Ocean acidification (continued)

Solubility of calcium carbonate

These may not seem like large changes in pH, but
 pH is a logarithmic scale.

Decrease by 0.1 pH units = increase in $[H^+]$ by a factor of 1.26

Q Causes a large relative change in CaCO₃ solubility:



• 13% increase in CaCO₃ solubility from pre-industrial CO₂ levels to now

Problem for shellfish!

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Vapor pressure of a pure substance

- Imagine taking a pure substance (solid or liquid), putting it in an air-tight container, removing all the air, then letting it come to equilibrium.
- The equilibrium pressure reached in this experiment is the vapor pressure of the substance and is due to evaporation of the substance.
 - Note: It isn't really necessary to remove the air, but it makes the measurement easier.
- Also note that if the partial pressure of a substance is lower than the (equilibrium) vapor pressure, it will evaporate. If, on the other hand, the partial pressure is higher than the vapor pressure, it will condense.

Problem: Calculate the vapor pressure of pure water at 25°C.

Species	$\Delta_f G^\circ/{ m kJ}{ m mol}^{-1}$	
$H_2O_{(I)}$	-237.140	
$H_2O_{(g)}$	-228.582	

Answer: 3.17×10^{-2} bar

Vapor pressures of solutions

- What about the vapor pressure of a solution?
- This could be due either to the solvent alone if the solute is (relatively) involatile (e.g. ionic compounds, sugar), or to a combination of the solvent and solute.

Vapor pressures of solutions Raoult's law

• Let's start by looking at the solvent, again using water as an example.

$$H_2O_{(I)} \rightleftharpoons H_2O_{(g)}$$

$$\mathcal{K} = rac{a_{ extsf{g}}}{a_{ extsf{l}}} = rac{p_{ extsf{H}_2 extsf{O}}/p^\circ}{X_{ extsf{H}_2 extsf{O}}}$$

• Note that $K = p(H_2O, pure)/p^\circ$. Notation: $p(H_2O, pure) \equiv p_{H_2O}^{\bullet}$

$$\therefore p_{\mathsf{H}_2\mathsf{O}} = p_{\mathsf{H}_2\mathsf{O}}^{\bullet} X_{\mathsf{H}_2\mathsf{O}}$$

or, in general,

$$p_{\text{solvent}} = p_{\text{solvent}}^{\bullet} X_{\text{solvent}}$$

• This equation is known as Raoult's law.

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- Let's calculate the vapor pressure of an aqueous solution prepared by dissolving 1.875 mol of sodium sulfate in 1.000 kg of water at 60°C.
- The vapor pressure of pure water at this temperature is 149 Torr.

$$n_{H_2O} = \frac{1000 \text{ g}}{18.0153 \text{ g/mol}} = 55.51 \text{ mol}$$

$$n_{Na^+} = 2(1.875 \text{ mol}) = 3.750 \text{ mol}$$

$$n_{SO_4^{2-}} = 1.875 \text{ mol}$$

$$\therefore X_{H_2O} = \frac{n_{H_2O}}{n_{H_2O} + n_{Na^+} + n_{SO_4^{2-}}} = \frac{55.51 \text{ mol}}{55.51 + 3.750 + 1.875 \text{ mol}}$$

$$= 0.9080$$

$$\therefore p_{H_2O} = p_{H_2O}^{\bullet} X_{H_2O} = (149 \text{ Torr})(0.9080) = 135 \text{ Torr}$$

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Vapor pressures of solutions

Henry's law

- We can also consider the vapor pressure of a volatile solute *or* the solubility of a gas in a solvent using equilibrium theory.
- For either of these cases, we consider

$$A_{(g)} \rightleftharpoons A_{(sol)}$$

A1 / O

$$\mathcal{K} = \frac{a(A, \text{sol})}{a(A, g)} = \frac{[A]/c^{\circ}}{p_A/p^{\circ}}$$

$$[A] = \frac{Kc^{\circ}}{p^{\circ}}p_A$$

or

- $[A] = k_H p_A$
- This equation is known as Henry's law, and the constant k_H is called the Henry's law constant.

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Vapor pressure

- Let us calculate the Henry's law constant for oxygen in water.
- $\bullet\,$ The standard free energy of formation of an aqueous oxygen molecule is $16.35\,kJ/mol.$
- The process is

$$O_{2(g)} \rightleftharpoons O_{2(aq)}$$

$$\begin{aligned} \Delta_r G^\circ &= \Delta_f G^\circ(\mathsf{O}_2, \mathsf{aq}) - \Delta_f G^\circ(\mathsf{O}_2, \mathsf{g}) = 16.35 \, \mathsf{kJ/mo} \\ \therefore \, \mathcal{K} &= \exp\left(-\frac{\Delta_r G^\circ_m}{RT}\right) \\ &= \exp\left(-\frac{16.35 \times 10^3 \, \mathsf{J/mol}}{(8.314 \, 463 \, \mathsf{J} \, \mathsf{K}^{-1} \mathsf{mol}^{-1})(298.15 \, \mathsf{K})}\right) \\ &= 1.366 \times 10^{-3} \\ k_H &= \frac{\mathcal{K}c^\circ}{p^\circ} = 1.366 \times 10^{-3} \, \mathsf{mol} \, \mathsf{L}^{-1} \mathsf{bar}^{-1} \end{aligned}$$

- Suppose that the atmospheric pressure of oxygen is 0.18 bar (roughly the case in Lethbridge, 920 m above sea level).
- Then, in water at equilibrium with the atmosphere, we have

$$\begin{split} [\mathsf{O}_2] &= k_H \, p_{\mathsf{O}_2} \\ &= \left(1.366 \times 10^{-3} \, \mathsf{mol} \, \mathsf{L}^{-1} \mathsf{bar}^{-1} \right) \left(0.18 \, \mathsf{bar} \right) \\ &= 2.5 \times 10^{-4} \, \mathsf{mol} / \mathsf{L}. \end{split}$$

 That's (roughly) how much oxygen the fishies in the Oldman River have to live on.